

# Huyber's Model of Glacial Cycles

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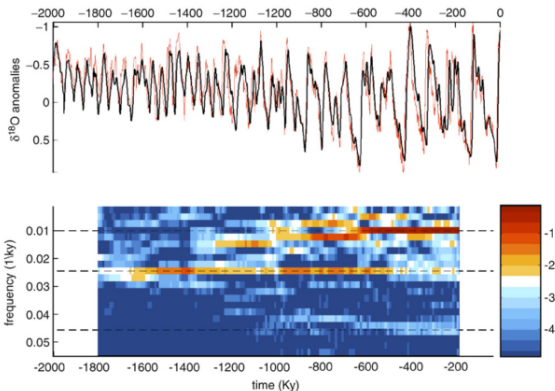
March 25, 2014

## Huybers' Model of Glacial Cycles

Huybers, P. Glacial variability over the last two million years: an extended depth-derived agemodel, continuous obliquity pacing, and the Pleistocene progression. *Quaternary Science Reviews*. 2007.

# The Obliquity Hypothesis

- $H_0$ : Deglaciations are independent of obliquity.
- $H_1$ : Deglaciations are triggered at a particular phase of Earth's obliquity.



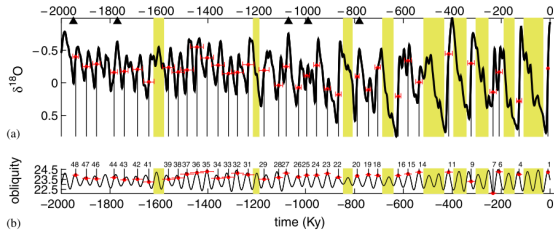
$\delta^{18}\text{O}$  record of last 2 Ma

## Rayleigh's $R$

- $\phi_n$  is the phase of obliquity sampled at the  $n^{\text{th}}$  deglacial event.
- Rayleigh's  $R$

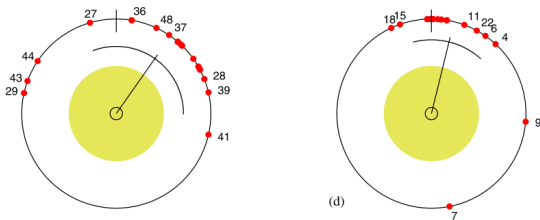
$$R = \frac{1}{N} \left| \sum_{n=1}^N \cos \phi_n + i \sin \phi_n \right|$$

# Deglacial events



Deglacial events and corresponding obliquity cycle

# Rayleigh circles

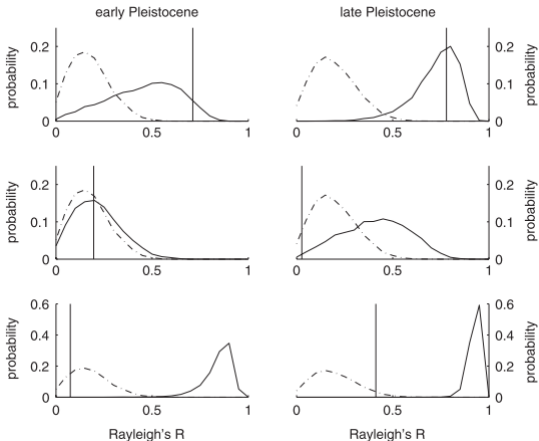


Rayleigh circles with obliquity phases plotted for early Pleistocene (left) and late Pleistocene (right)

## Rayleigh $R$ vs. Fourier analysis

- Nonlinearities associated with the variable duration and asymmetry do not affect the statistic.
- Errors in age model cause only linear changes in the phase, but distort the Fourier spectrum
- Requires fewer data points for statistical significance

# Rayleigh Statistic Results



Top: obliquity, middle: precession, bottom: eccentricity. The vertical line indicates the  $R$  value. Dashed lines indicate the probability distribution for  $H_0$ ; solid lines give the probability distribution for  $H_1$ .



# Rayleigh Statistic Results

	Early Pleistocene (2–1 Ma)				Late Pleistocene (1–0 Ma)			
	$R$	cv 1%	Power	Phase	$R$	cv 1%	Power	Phase
Obliquity	0.7	0.5	0.6	$\pm 56^\circ$	0.8	0.5	1.0	$\pm 28^\circ$
Precession	0.2	0.5	0.0	$\pm 88^\circ$	0.0	0.5	0.3	$\pm 56^\circ$
Eccentricity	0.1	0.5	1.0	$\pm 24^\circ$	0.4	0.5	1.0	$\pm 12^\circ$

Null hypothesis is rejected only for obliquity, implying that deglaciations are triggered at a particular phase of obliquity.

- The late Pleistocene deglacial events have  $R = 0.8$  vs.  $R = 0.7$  in the early Pleistocene.
- The late Pleistocene 100 Ka world has *greater* obliquity phase stability than the early Pleistocene 40 Ka world.

## Huybers' Model

$$\begin{aligned} V_t &= V_{t-1} + \eta_t \quad \text{and if } V_t \geq T_t \text{ terminate} \\ T_t &= at + b - c\theta'_t \end{aligned} \quad (1)$$

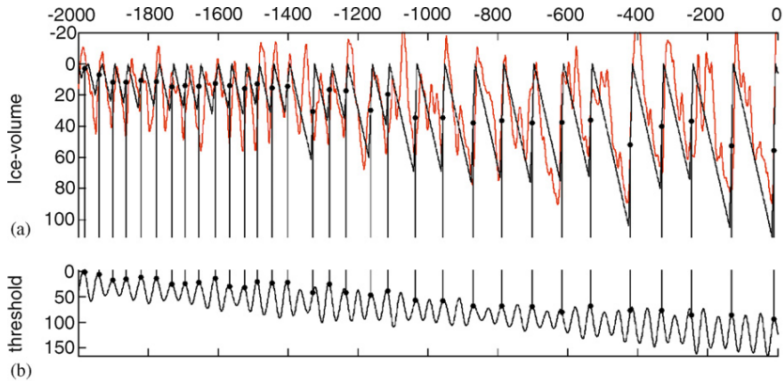
Upon termination, linearly reset  $V$  to 0 over 10 Ka

$V$  : ice volume

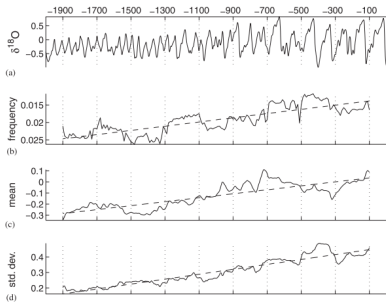
$T$  : deglaciation threshold

$\theta'$  : scaled obliquity

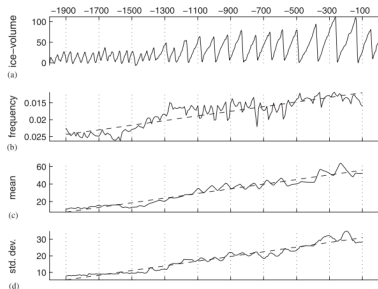
$\eta$  : ice volume growth rate



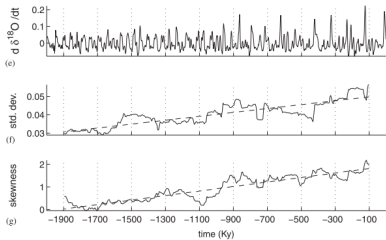
A deterministic run of the model



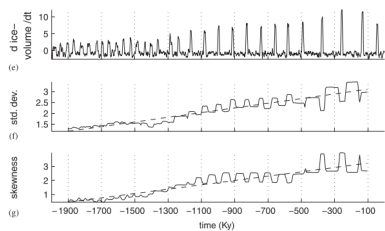
$\delta\text{O}^{18}$  data



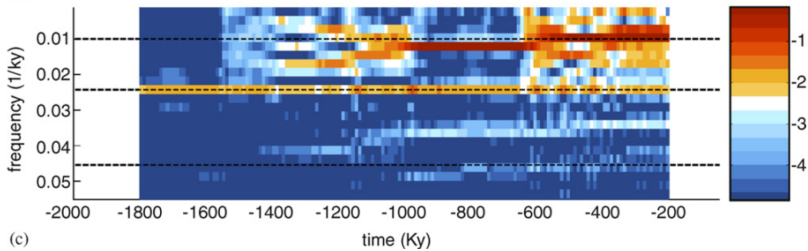
Modeled data



$\delta O^{18}$  data



Modeled data



Fourier transform of stochastic model

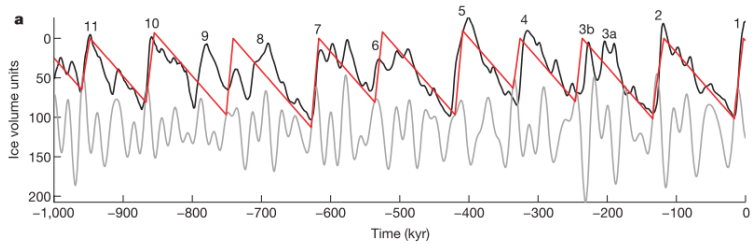
## Deglaciation Model with Combined Forcing

Huybers, P. Combined obliquity and precession pacing of late Pleistocene deglaciations. *Nature*. 2011.

$$V_t = V_{t-1} + \eta_t \quad \text{and if } V_t \geq T_t \text{ terminate}$$

$$T_t = 110 - 25\mathcal{F}_t$$

$$\mathcal{F}_t = \alpha^{1/2} e_t \sin(\omega_t - \phi) + (1 - \alpha)^{1/2} \epsilon_t$$



Deterministic model with combined forcing



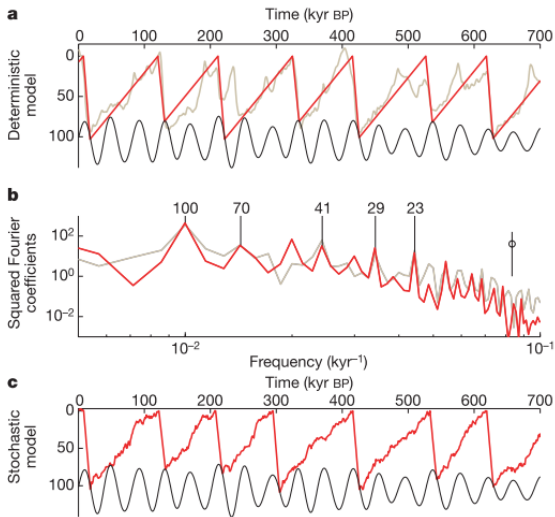
## Conclusion from 2011 paper

"Precession will tend to influence the precise timing of a deglaciation cycle, but obliquity will more fundamentally govern the interval between glaciations."

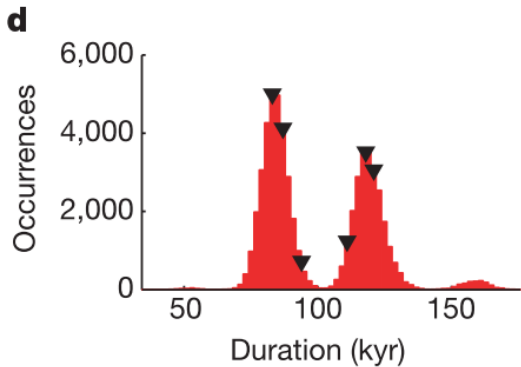
## Huybers and Wunsch original paper

Huybers, P. and Wunsch, C. Obliquity pacing of the late Pleistocene glacial terminations. *Nature*. 2005.

$$\begin{aligned} V_t &= V_{t-1} + \eta_t && \text{and if } V_t \geq T_t \text{ terminate} \\ T_t &= 100 - \theta'_t \end{aligned}$$



Deterministic and stochastic models with obliquity forcing



Histogram of time between terminations for many runs of stochastic model

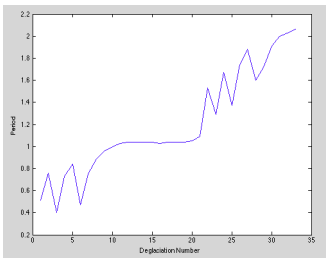
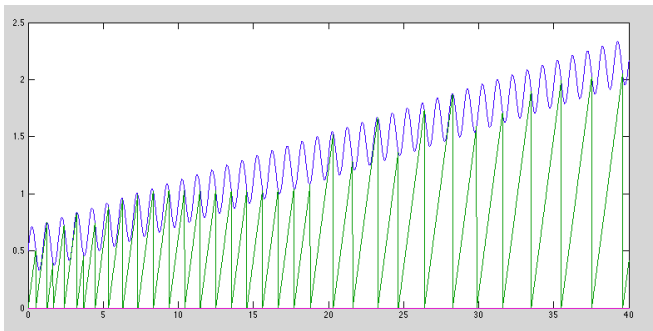
## Mathematical Questions

What happens in an idealized case when the forcing is just a sine curve?

$$V_t = V_{t-1} + \eta_t \quad \text{and if } V_t \geq T_t \text{ terminate}$$

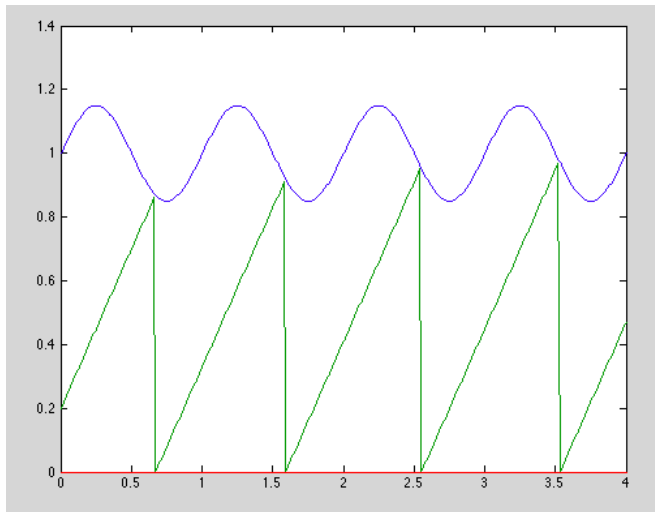
$$T_t = at + b + c \sin(2\pi t)$$

## Matlab simulation with sinusoidal obliquity



Progression of the  
period

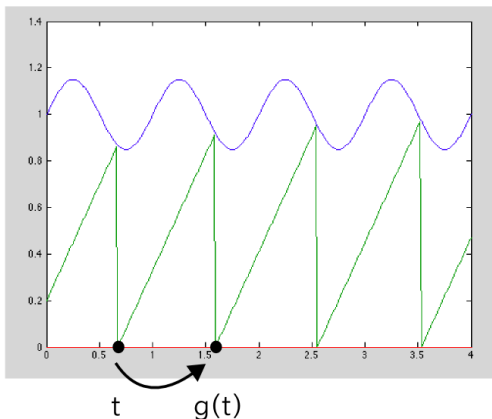
Simpler case: no increase over time



Suppose the threshold  $T(t)$  is periodic with period 1:

$$T(t+1) = T(t)$$

Let  $g : \mathbb{R} \rightarrow \mathbb{R}$  be the map taking a termination time  $t$  to the next termination time:





Suppose we start at a termination time  $x$ .

The next termination is at

$$g(x) = y \quad \text{where } y - x = T(y).$$

Then

$$T(y + 1) = T(y) = y - x = (y + 1) - (x + 1)$$

So  $y + 1$  would be next termination starting from time  $x + 1$ :

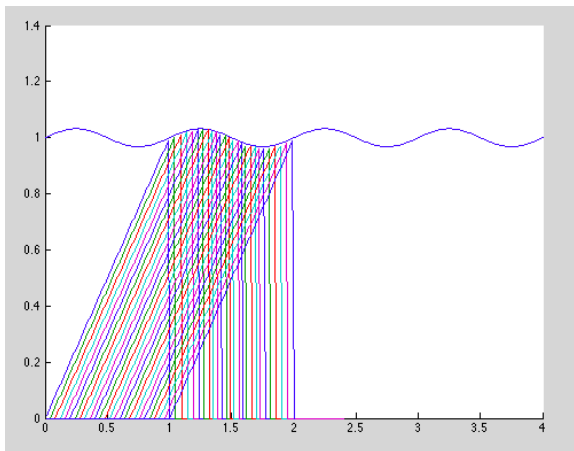
$$g(x + 1) = y + 1 = g(x) + 1$$

This means that

$$g(x + 1) = g(x) \pmod{1}$$

We can treat this simple model as a circle map  $f : [0, 1) \rightarrow [0, 1)$ :

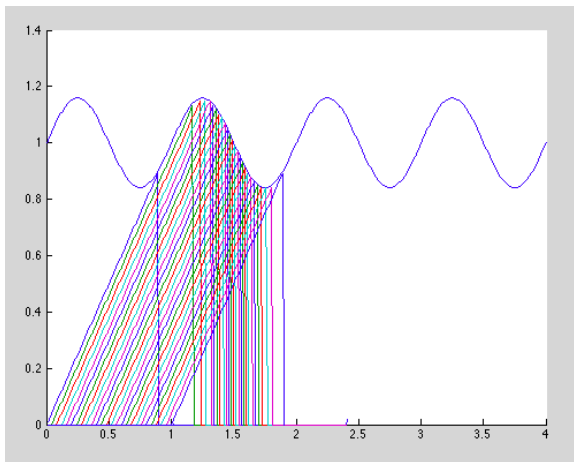
$$f(x) = y \pmod{1} \quad \text{where } y - x = A \sin(y) + B = T(y)$$



Smooth

We can treat this simple model as a circle map  $f : [0, 1) \rightarrow [0, 1)$ :

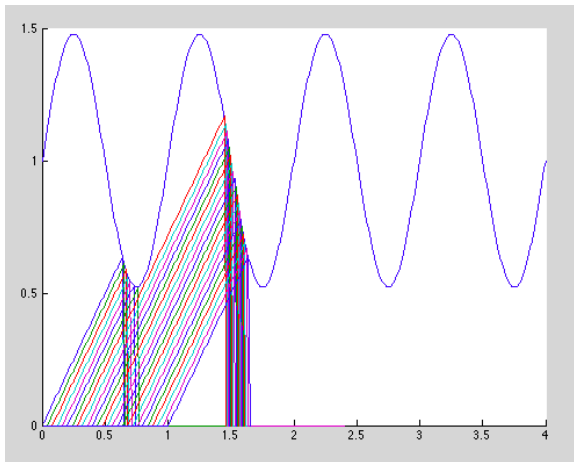
$$f(x) = y \pmod{1} \quad \text{where } y - x = A \sin(y) + B = T(y)$$



Continuous

We can treat this simple model as a circle map  $f : [0, 1) \rightarrow [0, 1)$ :

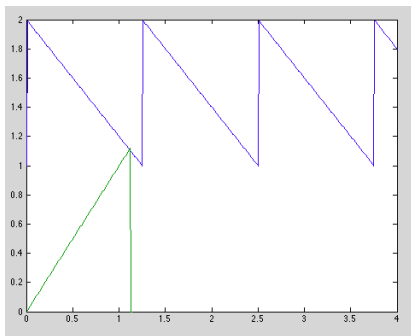
$$f(x) = y \pmod{1} \quad \text{where } y - x = A \sin(y) + B = T(y)$$



Discontinuous

## Even simpler circle maps

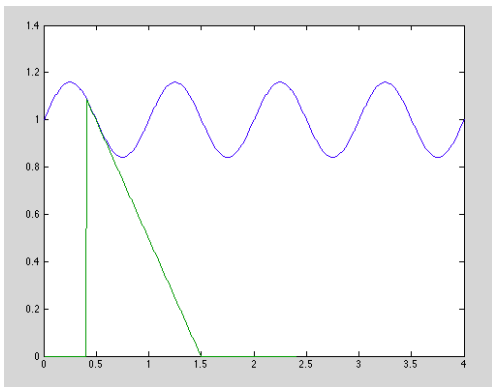
$$f(x) = mx + b \pmod{1}$$



- Canonical translation if  $m = 1$
- Not surjective if  $m < 1$
- Not injective if  $m > 1$

## Standard family of circle maps

$$f(x) = x + b + \frac{\omega}{2\pi} \sin(2\pi x) \pmod{1}$$



# Questions

- What is  $\lim_{n \rightarrow \infty} f^n(x)$  or  $\lim_{n \rightarrow \infty} f^n([0, 1])$ ?
- Is it possible to classify these maps by rotation numbers?
- Relation to Arnold tongues?
- Can we describe the transitions by increasing the threshold in Huybers' model?